

Weather Forecasting for Weather Derivatives

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We take a simple time series approach to modeling and forecasting daily average temperature in U.S. cities, and we inquire systematically as to whether it may prove useful from the vantage point of participants in the weather derivatives market. The answer is, perhaps surprisingly, yes. Time series modeling reveals conditional mean dynamics and, crucially, strong conditional variance dynamics in daily average temperature, and it reveals sharp differences between the distribution of temperature and the distribution of temperature surprises. As we argue, it also holds promise for producing the long-horizon predictive densities crucial for pricing weather derivatives, so that additional inquiry into time series weather forecasting methods will likely prove useful in weather derivatives contexts.

KEY WORDS: Financial derivatives; Hedging; Insurance; Risk management; Seasonality; Temperature.

1. INTRODUCTION

Weather derivatives are a fascinating new type of security, making prespecified payouts if prespecified weather events occur. The market has grown rapidly. In 1997, the market for weather derivatives was nonexistent. In 1998, the market was estimated at \$500 million, but it was still illiquid, with large spreads and limited secondary market activity. More recently, the market has grown to more than \$5 billion, with better liquidity. Outlets such as the *Weather Risk* (e.g., 1998, 2000) supplements to *Risk Magazine* have chronicled the development.

Weather derivative instruments include weather swaps, options, and option collars (see, e.g., Geman 1999; Dischel 2002 for definitions and descriptions). The payoffs of these instruments may be linked to various “underlying” weather-related variables, including heating degree days, cooling degree days, growing degree days, average temperature, maximum temperature, minimum temperature, precipitation (rainfall, snowfall), humidity, and sunshine, among others—even the National Weather Service’s temperature forecast for the coming week. Most trading is over the counter, but exchange-based trading is gaining momentum. For example, temperature-related derivatives, are traded on the Chicago Mercantile Exchange (CME) for major U.S. cities.

A number of interesting considerations make weather derivatives different from “standard” derivatives. First, the underlying object (weather) is not traded in a spot market. Second, unlike financial derivatives, which are useful for price hedging but not for quantity hedging, weather derivatives are useful for quantity hedging but not necessarily for price hedging (although the two are obviously related). That is, weather derivative products provide protection against weather-related changes in quantities, complementing extensive commodity price risk management tools already available through futures. Third, although liquidity in weather derivative markets has improved, it will likely

never be as good as in traditional commodity markets, because weather is by its nature a location-specific and nonstandardized commodity, unlike, say, a specific grade of crude oil.

Interestingly, weather derivatives are also different from insurance. First, there is no need to file a claim or prove damages. Second, there is little moral hazard. Third, unlike insurance, weather derivatives allow one to hedge against comparatively good weather in other locations, which may be bad for local business (e.g., a bumper crop of California oranges may lower the prices received by Florida growers).

Weather forecasting is crucial to both the demand and the supply sides of the weather derivatives market. Consider first the demand side; any firm exposed to weather risk on either the output (revenue) side or the input (cost) side is a candidate for productive use of weather derivatives. This includes obvious players, such as energy companies, utilities, and insurance companies, and less obvious players, such as ski resorts, grain millers, cities facing snow-removal costs, consumers who want fixed heating and air-conditioning bills, and firms seeking to avoid financial writedowns due to weather-driven poor performance. The mere fact that such agents face weather fluctuations, however, does not ensure a large hedging demand, because even very large weather fluctuations would create little weather risk if they were highly predictable. Weather risk, then, is about the *unpredictable* component of weather fluctuations—“weather surprises,” or “weather noise.” To assess the potential for hedging against weather surprises, and to formulate the appropriate hedging strategies, one needs to determine how much weather noise exists for weather derivatives to eliminate. This requires a weather model. What does weather noise look like over space and time? What is its distribution? Answering such questions requires statistical weather modeling and forecasting, the topic of this article.

Now consider the supply side: sellers of weather derivatives who want to price them, arbitrageurs who want to exploit situations of apparent mispricing, and so on. How should weather derivatives be priced? It seems clear that standard approaches to arbitrage-free pricing (e.g., Black and Scholes 1973) are inapplicable in weather derivative contexts. In particular, there is in general no way to construct a portfolio of financial assets that replicates the payoff of a weather derivative. Hence, the only way to price options reliably is by using forecasts of the underlying weather variable, in conjunction with a utility function, as argued by, for example, Davis (2001). This again raises the

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crucial issue of how to construct good weather forecasts (not only point forecasts, but also—and crucially—complete density forecasts), potentially at horizons much longer than those commonly emphasized by meteorologists. Hence, the supply-side questions, like the demand-side questions, are intimately related to weather modeling and forecasting.

Curiously, however, it seems that little thought has been given to the crucial question of how best to approach the weather modeling and forecasting that underlies weather derivative demand and supply. The meteorological weather forecasting literature focuses primarily on short-horizon point forecasts produced from structural physical models of atmospheric conditions (see, e.g., the overview in Tribia 1997). Although such an approach is best for helping one decide how warmly to dress tomorrow, it is not at all obvious that it is best for producing the long-horizon density forecasts relevant for weather derivatives. In particular, successful forecasting does not necessarily require a structural model; over the last 30 years statisticians and econometricians have made great strides in using nonstructural models of time series trend, seasonal, and cyclical components to produce good forecasts, including long-horizon density forecasts (for a broad overview, see Diebold 2004).

In this article, then, motivated by considerations related to the weather derivatives market, we take a nonstructural time series approach to temperature modeling and forecasting, systematically asking whether it proves useful. We are not the first to adopt a time series approach, although the literature is sparse and inadequate for our purposes. The analyses of Harvey (1989), Hyndman and Grunwald (2000), Milionis and Davies (1994), Visser and Molenaar (1995), Jones (1996), and Pozo, Esteban-Parra, Rodrigo, and Castro-Diez (1998) suggest its value, for example, but they do not address the intrayear temperature forecasting relevant to our concerns. Seater (1993) studied long-run temperature trend, but little else. Contemporaneous and independent work by Cao and Wei (2001) and Torro, Meneu, and Valor (2001) considered time series models of average temperature, but their models are more restrictive and their analyses more limited.

We contribute by providing insight into both conditional mean dynamics and conditional variance dynamics of daily average temperature, as relevant for weather derivatives. Strong conditional variance dynamics are a central part of the story. We also highlight the differences between the distributions of weather and weather *innovations*. Finally, we evaluate the performance of time series point and density forecasts as relevant for weather derivatives. The results are mixed but ultimately encouraging, and they point toward directions that may yield future forecasting improvements.

We proceed as follows. In Section 2 we discuss our data and our focus on modeling and forecasting daily average temperature, and we report the results of time series modeling. In Section 3 we report the results of out-of-sample point and density forecasting exercises. In Section 4 we offer concluding remarks and highlight some pressing directions for future research.

2. TIME SERIES WEATHER DATA AND MODELING

Here we discuss our choice of weather data and its collection. We are interested in daily average temperature (T), which is widely reported and followed. Moreover, the heating degree days (HDDs) and cooling degree days (CDDs) on which weather derivatives are commonly written are simple transformations of daily average temperature. We directly model and forecast daily average temperature, measured in degrees Fahrenheit, for each of four measurement stations (Atlanta, Chicago, Las Vegas and Philadelphia) for 1/1/60 through 11/05/01, resulting in 15,285 observations per measurement station. Each of the cities is one of the 10 for which temperature-related weather derivatives are traded at the CME. In earlier and longer versions of this article (Campbell and Diebold 2002, 2003), we report results for all 10 cities, which are qualitatively identical. We obtained the data from Earth Satellite (EarthSat) corporation; they are precisely those used to settle temperature-related weather derivative products traded on the CME. The primary underlying data source is the National Climatic Data Center (NCDC), a division of the National Oceanographic and Atmospheric Administration. Each of the measurement stations supplies its data to the NCDC, and those data are in turn collected by EarthSat.

Before proceeding to detailed modeling and forecasting results, it is useful to get an overall feel for the daily average temperature data. Figure 1 plots the daily average temperature series for the last 5 years of the sample. The time series plots reveal strong and unsurprising seasonality in average temperature; in each city, the daily average temperature moves repeatedly and regularly through periods of high temperature (summer) and low temperature (winter). Importantly, however, the seasonal fluctuations differ noticeably across cities in terms of both amplitude and detail of pattern.

Figure 2 shows how the seasonality in daily average temperature manifests itself in unconditional temperature densities. The densities are either bimodal or nearly so, with peaks characterized by cool and warm temperatures. Also, with the exception of Las Vegas, each density is negatively skewed. The distributional results are in line with those of von Storch and Zwiers (1999), who noted that although daily average temperature of

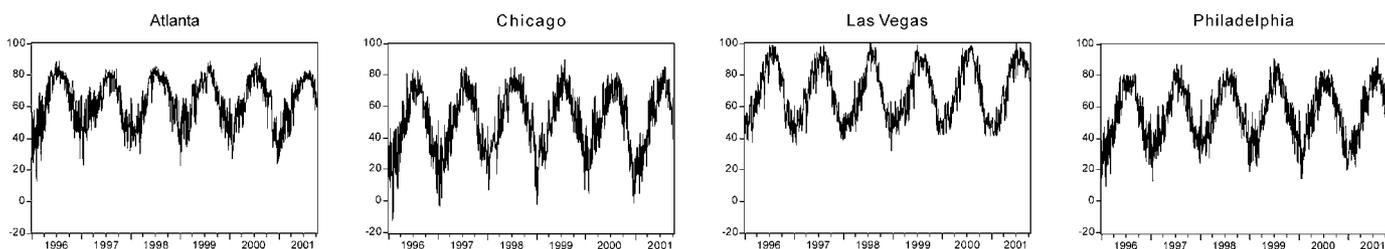


Figure 1. Time Series Plots, Daily Average Temperature. Each panel displays a time series plot of daily average temperature, 1996–2001.

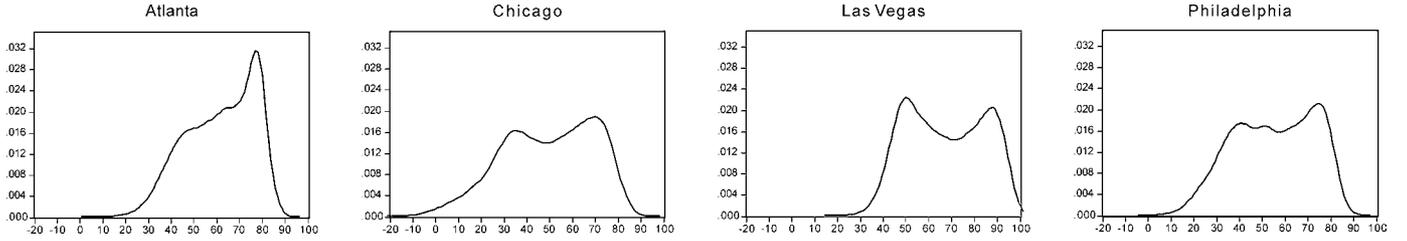


Figure 2. Estimated Unconditional Distributions, Daily Average Temperature. Each panel displays a kernel density estimate of the unconditional distribution of daily average temperature, 1960–2001. In each case we use the Epanechnikov kernel and select the bandwidth using Silverman's rule, $h = .9\hat{\sigma} N^{-.2}$.

ten appears Gaussian if studied over sufficiently long times in the troposphere, daily average surface temperatures may have different distributions, and with Neese (1994), who documented skewness and bimodality in daily maximum temperatures.

The discussion thus far suggests that a seasonal component will be important in any time series model fit to daily average temperature, because average temperature displays pronounced seasonal variation, with the seasonal patterns differing noticeably across cities. We use a low-ordered Fourier series to model this seasonality, the benefits of which are twofold. First, it produces a smooth seasonal pattern, which accords with the basic intuition that the progression through different seasons is gradual rather than discontinuous. Second, it promotes parsimony, which enhances numerical stability in estimation. Such considerations are of relevance given the rather large size of our dataset (roughly 15,000 daily observations for each of four cities) and the numerical optimization that we subsequently perform.

One naturally suspects that nonseasonal factors may also be operative in the dynamics of daily average temperature. One such factor is trend, which may be relevant but is likely minor, given the short 40 year span of our data. We therefore simply allow for a simple low-ordered polynomial deterministic trend. Another such factor is cycle, by which we mean any sort of persistent covariance stationary dynamics apart from seasonality and trend. We capture cyclical dynamics using autoregressive lags.

The discussion thus far has focused on conditional mean dynamics, with contributions coming from trend, seasonal, and cyclical components. We also allow for conditional variance (volatility) dynamics, with contributions coming from both seasonal and cyclical components. We approximate the seasonal volatility component using a Fourier series, and we approximate the cyclical volatility component using a generalized autoregressive conditional heteroscedasticity (GARCH) process (Engle 1982; Bollerslev 1986).

Assembling the various pieces, we estimate the following daily average temperature model for each of our four cities:

$$T_t = \text{Trend}_t + \text{Seasonal}_t + \sum_{l=1}^L \rho_{t-l} T_{t-l} + \sigma_t \varepsilon_t, \quad (1)$$

where

$$\text{Trend}_t = \sum_{m=0}^M \beta_m t^m, \quad (1a)$$

$$\text{Seasonal}_t = \sum_{p=1}^P \left(\sigma_{c,p} \cos\left(2\pi p \frac{d(t)}{365}\right) + \sigma_{s,p} \sin\left(2\pi p \frac{d(t)}{365}\right) \right), \quad (1b)$$

$$\sigma_t^2 = \sum_{q=1}^Q \left(\gamma_{c,q} \cos\left(2\pi q \frac{d(t)}{365}\right) + \gamma_{s,q} \sin\left(2\pi q \frac{d(t)}{365}\right) \right) + \sum_{r=1}^R \alpha_r (\sigma_{t-r} \varepsilon_{t-r})^2 + \sum_{s=1}^S \beta_s \sigma_{t-s}^2, \quad (1c)$$

$$\varepsilon_t \sim \text{iid}(0, 1), \quad (1d)$$

and $d(t)$ is a repeating step function that cycles through $1, \dots, 365$ (as we drop February 29 in all leap years). In all that follows, we set $L = 25$, $M = 1$, $P = 3$, $Q = 3$, $R = 1$, and $S = 1$, which both the Akaike and Schwarz information criteria indicate are more than adequate for each city. Maintaining the rather large value of $L = 25$ costs little given the large number of available degrees of freedom, and it helps capture long-memory dynamics, if present, as suggested by results such as those of Bloomfield (1992). Following Bollerslev and Wooldridge (1992), we consistently estimate this regression model with GARCH disturbances by Gaussian quasi-maximum likelihood.

Now let us discuss the estimation results, starting with the conditional-mean model [(1), (1a), and (1b)]. First, and perhaps surprisingly, most cities display a statistically significant trend in daily average temperature. In most cases, the trend is much larger than the increase in average global temperature over the same period. For example, the results indicate that the daily average temperature in Atlanta has increased by 3°F in the last 40 years. Such large trend increases are likely a consequence of development and air pollution that increased urban temperatures in general, and urban airport temperatures in particular, where most of the U.S. recording stations are located, a phenomenon often dubbed the “heat island effect.” Second, the conditional mean dynamics display both statistically significant and economically important seasonality. Third, conditional mean dynamics also display strong cyclical persistence. The estimated autoregressions display an interesting root pattern, common across all four cities, regardless of location. The dominant root is large and real, around .85; the second and third roots are a complex conjugate pair with moderate modulus, around .3; and all subsequent roots are much smaller in modulus.

Figure 3 shows kernel estimates of residual densities. Four features emerge. First, average temperature residuals are *much* less variable than average temperature itself; that is, weather

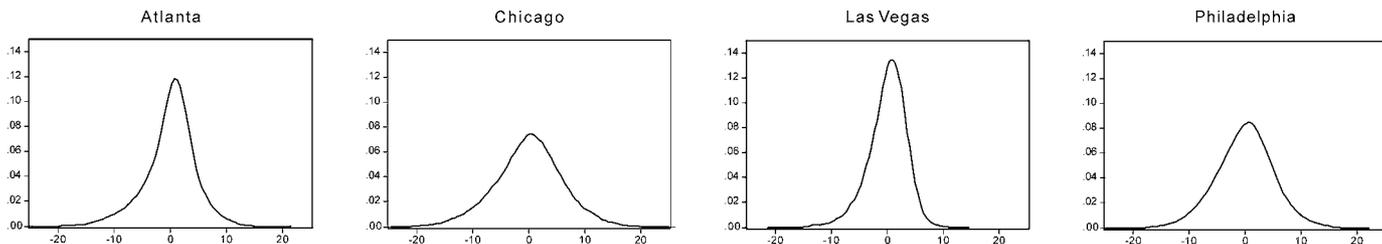


Figure 3. Estimated Unconditional Distribution, of Residuals, Daily Average Temperature. Each panel displays a kernel density estimate of the distribution of the residuals from our daily average temperature model, $T_t - \hat{Trend}_t - \hat{Seasonal}_t - \sum_{i=1}^{25} \hat{\rho}_i T_{t-i}$. In each case we use the Epanechnikov kernel and select the bandwidth using Silverman's rule, $h = .9\hat{\sigma} N^{-.2}$.

surprises are much less variable than the weather itself, with residual standard deviations only one-third or so of the average temperature standard deviations. Second, again as expected, all residual densities are unimodal as opposed to bimodal, with contrast to the unconditional densities examined earlier, due to the model's success in capturing seasonal highs and lows. Third, the spreads of the residual densities vary noticeably across cities, indicating that weather risk is much greater in some cities than in others. Fourth, all of the residual densities have only moderate negative skewness and moderate excess kurtosis; the average residual skewness and kurtosis coefficients are $-.36$ and 4.10 .

All told, the conditional mean model [(1), (1a), and (1b)] fits quite well, with R^2 typically above 90%. Figure 4 shows model residuals ($\hat{\sigma}_t \hat{\varepsilon}_t$) over the last 5 years of the estimation sample, which provide a first glimpse of an important phenomenon: pronounced and persistent time-series volatility dynamics (conditional heteroscedasticity) in the temperature shocks. In particular, weather risk, as measured by its innovation variance, appears to be seasonal, as the amplitude of the residual fluctuations varies over the course of each year, widening in winter and narrowing in summer. It seems that such seasonal conditional heteroscedasticity in temperature was first noted, informally, in an economic context by Roll (1984). Our volatility model (1c) is designed to approximate the conditional heteroscedasticity formally and flexibly.

To gain additional insight into the strength and pattern of the conditional heteroscedasticity, Figure 5 displays correlograms of the squared residuals, taken to a maximum displacement of 800 days. There is clear evidence of strong nonlinear residual dependence, driven by strong conditional variance dynamics. This contrasts sharply with the correlograms of the residuals (not shown, to conserve space), which are indicative of weak white noise.

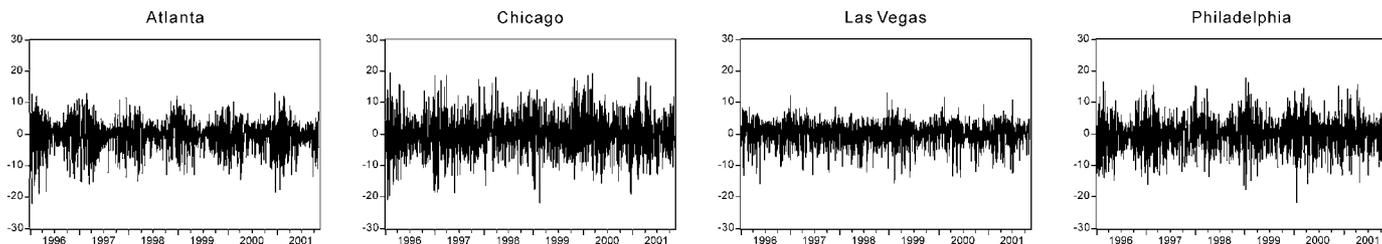


Figure 4. Estimated Model Residuals, Daily Average Temperature. Each panel displays the residuals from an unobserved-components model, $T_t = Trend_t + Seasonal_t + \sum_{i=1}^{25} \rho_i T_{t-i} + \sigma_t \varepsilon_t$, for 1996–2001.

In keeping with the results of the correlogram analysis of squared residuals, and in parallel to the aforementioned results for the estimated conditional mean function, the estimated conditional variance function (1c) reveals both significant seasonal (Fourier) and cyclical (GARCH) components. The Fourier part appears to capture adequately all volatility seasonality, whereas the GARCH part captures the remaining nonseasonal volatility persistence. The seasonal volatility component is the relatively more important; it is significant and sizeable for all cities. The nonseasonal GARCH volatility component has a smaller effect, and there is considerable range in the estimates of nonseasonal volatility persistence, as determined by the GARCH parameter, β , implying different half-lives of nonseasonal volatility shocks across cities. For example, the half-life of a Las Vegas nonseasonal volatility shock is approximately 1 day, whereas the half-life of a Chicago nonseasonal volatility shock is approximately 7 days.

Figure 6 plots the estimated residual conditional standard deviation from 1996 through 2001. The basic pattern is one of strong seasonal volatility variation, with additional GARCH volatility effects, the persistence of which varies across cities. For each city, seasonal volatility appears to be highest during the winter months. Among other things, this indicates that correct pricing of weather derivatives may in general be crucially dependent on the season covered by the contract. Some cities display a great deal of seasonal volatility variation; the conditional standard deviation of Atlanta temperature shocks, for example, roughly triples each winter, whereas seasonal volatility varies less in other cities, such as Las Vegas.

It is interesting to note from Figure 6 that the GARCH volatility effects appear more pronounced in winter, when the volatility seasonal component is high, which might suggest the desirability of a multiplicative volatility specification. Nelson's (1991) exponential GARCH(1, 1) is one such attractive specification, replacing the volatility specification for σ_t^2 in (1c) with

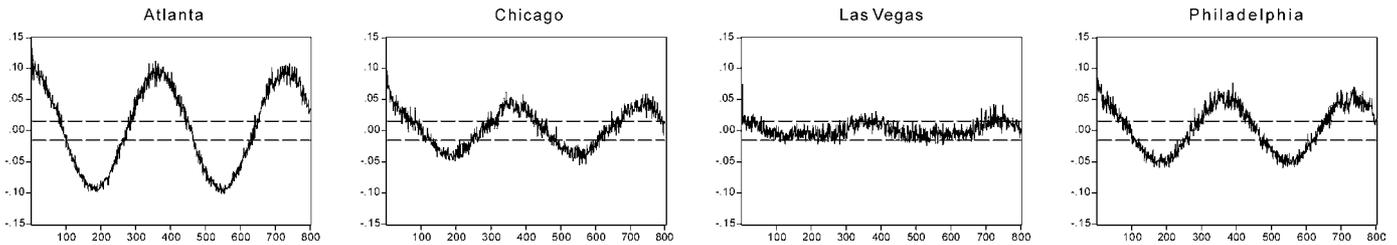


Figure 5. Correlogram of Squared Residuals, Daily Average Temperature. Each panel displays sample autocorrelations of the squared residuals from our daily average temperature model, $(T_t - \widehat{Trend}_t - \widehat{Seasonal}_t - \sum_{i=1}^{25} \hat{\rho}_i T_{t-i})^2$, together with Bartlett's approximate 95% confidence intervals under the null hypothesis of white noise.

an alternative but related specification for $\ln \sigma_t^2$. Estimation of exponential GARCH models produced fitted conditional variance series nearly identical to those of the original GARCH models, however.

We also estimated the densities of the standardized residuals, $(T_t - \widehat{T}_t)/\hat{\sigma}_t$, where \widehat{T}_t is the fitted value of daily average temperature (not shown, to conserve space). They still display negative skewness, and the average across cities is $-.45$. Modeling the conditional heteroscedasticity does, however, reduce (but not completely eliminate) residual excess kurtosis; the average across cities is now 3.74. Finally, we also computed the correlograms of squared standardized residuals (again not shown, to conserve space); there was no significant deviation from white noise behavior, indicating that the fitted model (1) is adequate.

3. TIME SERIES WEATHER FORECASTING

Armed with a hopefully adequate time series model for daily average temperature, we now proceed to examine its performance in out-of-sample weather forecasting. We begin by examining its performance in short-horizon point forecasting, despite the fact that short horizons and point forecasts are not of maximal relevance for weather derivatives, to compare our performance to that of a very sophisticated leading meteorological forecast. One naturally suspects that the much larger information set on which the meteorological forecast is based will result in superior short-horizon point forecasting performance, but even if this is so, of great interest is the question of how quickly and with what pattern the superiority of the meteorological forecast deteriorates with forecast horizon.

We then progress to assess the performance of our model's long-horizon density forecasts, which are of maximal interest in weather derivative contexts, given the underlying option pricing considerations, and which let us explore the effects of using

a daily model to produce much longer-horizon density forecasts. Simultaneously, we also move to forecasting HDD_t rather than T_t , which lets us match the most common weather derivative "underlying."

3.1 Point Forecasting

We assess the short-term accuracy of daily average temperature forecasts based on our seasonal + trend + cycle model. In what follows we refer to those forecasts as "autoregressive forecasts," for obvious reasons. We evaluate the autoregressive forecasts relative to three benchmark competitors, ranging from rather naive to very sophisticated. The first benchmark forecast is a no-change forecast. The no-change forecast, often called the "persistence forecast" in the climatological literature, is the minimum mean squared error forecast at all horizons if daily average temperature follows a random walk.

The second benchmark forecast is from a more sophisticated two-component (seasonal + trend) model. It captures (daily) seasonal effects via day-of-year dummy variables, in keeping with the common climatological use of daily averages as benchmarks, and captures trend via a simple linear deterministic function of time. We refer to this forecast as the "climatological forecast."

The third benchmark forecast, unlike benchmarks one and two, is not at all naive; on the contrary, it is a highly sophisticated forecast produced in real time by EarthSat. To produce their forecast, EarthSat meteorologists pool their expert judgement with model-based numerical weather prediction (NWP) forecasts from the National Weather Service, as well as with forecasts from European, Canadian, and U.S. Navy weather services. This blending of judgement with models is typical of best-practice modern weather forecasting.

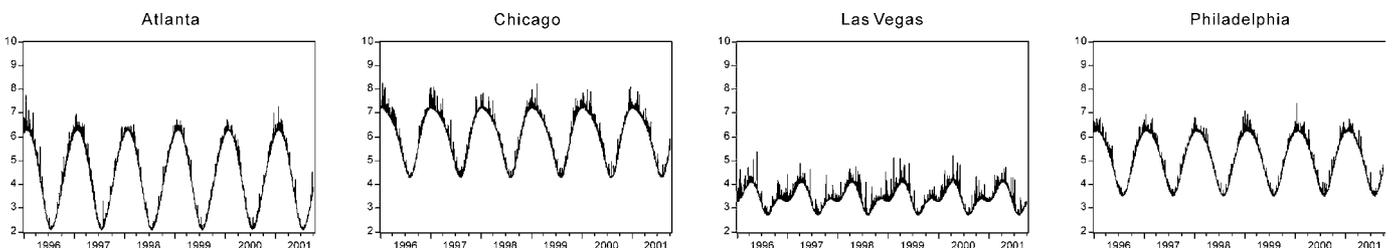


Figure 6. Estimated Conditional Standard Deviations, Daily Average Temperature. Each panel displays a time series of estimated conditional standard deviations ($\hat{\sigma}_t$) of daily average temperature, where $\hat{\sigma}_t^2 = \sum_{q=1}^3 (\hat{\gamma}_{c,q} \cos(2\pi q \frac{\pi(t)}{365}) + \hat{\gamma}_{s,q} \sin(2\pi q \frac{\pi(t)}{365})) + \hat{\alpha}(\sigma_{t-1}\varepsilon_{t-1})^2 + \hat{\beta}\hat{\sigma}_{t-1}^2$, for 1996–2001.

Table 1. Point Forecast Accuracy Comparisons Daily Average Temperature

	1-day-ahead	3-day-ahead	5-day-ahead	7-day-ahead	9-day-ahead	11-day-ahead
Atlanta						
Persistence	4.50	8.00	8.72	9.07	8.99	9.28
Climatological	6.93	6.88	6.84	7.04	6.93	6.59
Autoregressive	4.12	6.45	6.69	7.03	6.89	6.59
EarthSat	2.74	3.84	5.10	6.04	6.65	7.00
Chicago						
Persistence	6.73	10.50	11.06	11.54	11.74	11.99
Climatological	8.74	8.72	8.72	8.50	8.88	8.55
Autoregressive	6.06	8.38	8.57	8.45	8.84	8.53
EarthSat	3.22	4.70	6.31	7.46	8.48	8.92
Las Vegas						
Persistence	3.78	6.15	7.08	7.71	7.96	7.93
Climatological	5.99	5.85	5.80	6.02	5.97	5.80
Autoregressive	3.57	5.20	5.58	5.92	5.97	5.78
EarthSat	2.54	3.28	4.19	5.32	5.81	6.04
Philadelphia						
Persistence	5.53	8.83	9.83	9.87	9.55	10.18
Climatological	7.12	6.95	7.27	7.19	6.98	7.19
Autoregressive	4.95	6.74	7.23	7.15	6.95	7.08
EarthSat	2.61	3.91	5.35	6.26	7.24	8.37

NOTE: Each forecast's RMSE, measured in degrees Fahrenheit.

We were able to purchase approximately 2 years of forecasts from EarthSat. The sample period runs from 10/11/99 (the date when EarthSat began to archive their forecasts electronically and make them publicly available) through 10/22/01. Each weekday, EarthSat makes a set of h -day-ahead daily average temperature forecasts, for $h = 1, 2, \dots, 11$. EarthSat does not make forecasts on weekends.

We measure accuracy of all point forecasts using h -day-ahead root mean squared prediction error (RMSPE). We assess point forecasting accuracy at horizons of $h = 1, 2, \dots, 11$ days, because those are the horizons at which EarthSat's forecasts are available. We compute measures of the accuracy of our model and the EarthSat model relative to that of the persistence and climatological benchmarks. RMSPE ratios relative to benchmarks are called *skill scores* in the meteorological literature (Brier and Allen 1951) and *U-statistics* in the econometrics literature (Theil 1966). Specifically, in an obvious notation, the skill score relative to the persistence forecast is $Skill_h^p = \sqrt{\sum(\hat{T}_{t+h,t} - T_{t+h})^2 / \sum(T_{t+h,t}^p - T_{t+h})^2}$, where $T_{t+h,t}^p = T_t$ is the persistence forecast and $\hat{T}_{t+h,t}$ is either the autoregressive forecast or the EarthSat forecast. The skill score relative to the climatological forecast is $Skill_h^c = \sqrt{\sum(\hat{T}_{t+h,t} - T_{t+h})^2 / \sum(\hat{T}_{t+h,t}^c - T_{t+h})^2}$, where $\hat{T}_{t+h,t}^c$ denotes

the climatological forecast, $\hat{T}_{t+h,t}^c = \hat{\beta}_0 + \hat{\beta}_1(t+h) + \sum_{i=1}^{365} \hat{\delta}_i D_{i,t+h}$, and D_{it} is a daily dummy.

A number of nuances merit discussion. First, for each of our time series models, we estimate and forecast recursively, using only the data available in real time. Thus at any time our forecasts use no more average temperature information than do EarthSat's. In fact, our forecasts are based on *less* average temperature information; our forecast for day $t+1$ made on day t is based on daily average temperature through 11:59 PM of day t , whereas the EarthSat forecast for day $t+1$, which is not released until 6:45 AM on day $t+1$, potentially makes use of the history of temperature through 6:45 AM of day $t+1$. Second, we make forecasts using our models only on the dates that EarthSat made forecasts. In particular, we make no forecasts on weekends. Hence our accuracy comparisons proceed by averaging squared errors over precisely the same days as those corresponding to the EarthSat errors. This ensures a fair comparison.

Table 1 reports RMSPEs at horizons of $h = 1, 3, 5, 7, 9$, and 11 days, for all cities and forecasting models. In addition, skill scores are graphed as a function of horizon, against the persistence forecast in Figure 7 and against the climatological forecast in Figure 8, for all cities and horizons. The results are the same for all cities, so it is not necessary to discuss them individually by city. The results most definitely do differ across

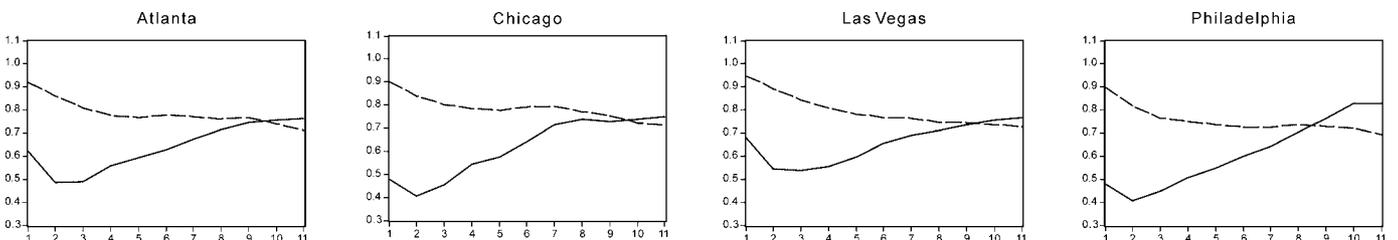


Figure 7. Forecast Skill Relative to Persistence Forecast, Daily Average Temperature Point Forecasts. Each panel displays the ratio of a forecast's RMSPE to that of a persistence forecast, for 1-day-ahead through 11-day-ahead horizons. The solid line represents the EarthSat forecast, and the dashed line represents the autoregressive forecast. The forecast evaluation period is 10/11/99–10/22/01.

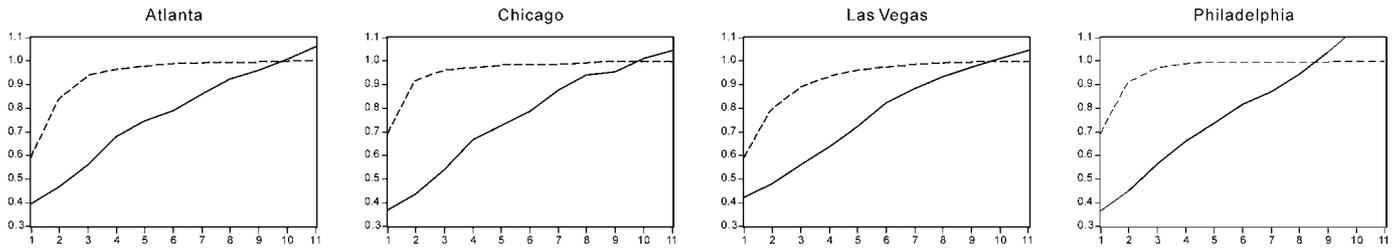


Figure 8. Forecast Skill Relative to Climatological Forecast, Daily Average Temperature. Each panel displays the ratio of a forecast's RMSPE to that of a climatological forecast, for 1-day-ahead through 11-day-ahead horizons. The solid line represents the EarthSat forecast, and the dashed line represents the autoregressive forecast. The forecast evaluation period is 10/11/99–10/22/01.

models and horizons, however, as we now discuss. We first discuss the performance of the time series forecasts, then discuss the EarthSat forecasts.

We first consider the forecasting performance of the persistence, climatological, and autoregressive models across the various horizons. First, consider the comparative performance of the persistence and climatological forecasts. When $h = 1$, the climatological forecasts are much worse than the persistence forecasts, reflecting the fact that persistence in daily average temperature renders the persistence forecast quite accurate at very short horizons. As the horizon lengthens, however, this result is reversed; the persistence forecast becomes comparatively poor, because the temperature today has rather little to do with the temperature, for example, 9 days from now.

Second, consider the performance of the autoregressive forecasts relative to the persistence and climatological forecasts. Even when $h = 1$, the autoregressive forecasts consistently outperform the persistence forecast, and the relative superiority of the autoregressive forecasts *increases* with horizon. The autoregressive forecasts also outperform the climatological forecasts at short horizons, but their comparative superiority *decreases* with horizon. The performance of the autoregressive forecast is commensurate with that of the climatological forecast roughly by the time $h = 4$, indicating that the cyclical dynamics captured by the autoregressive model via the inclusion of lagged dependent variables, which are responsible for its superior performance at shorter horizons, are not very persistent and thus are not readily exploited for superior forecast performance at longer horizons.

We now compare the forecasting performance of the autoregressive model and the EarthSat model. When $h = 1$, the EarthSat forecasts are much better than the autoregressive forecasts (which in turn are better than either the persistence forecast or the climatological forecast, as discussed earlier). Figures 7 and 8 make clear, however, that the EarthSat forecasts outperform the autoregressive forecasts by progressively less as the horizon lengthens, with nearly identical performance obtaining by the time $h = 8$. One could even make a case that the point forecasting performances of EarthSat and our three-component model become indistinguishable before $h = 8$ (say, by $h = 5$) if one were to account for the sampling error in the estimated RMSPEs and for the fact that the EarthSat information set for any day t actually contains a few hours of the next day.

Thus far we have examined our model's performance in short-horizon point forecasting, to compare it with competitors

such as EarthSat, who produce only short-horizon point forecasts. Its point forecasting performance is not particularly encouraging; although it appears no worse than its competitors at horizons of 8 or 10 days, it also appears no better. The nature of temperature dynamics simply makes *any* point forecast of temperature unlikely to beat the climatological forecast at long horizons, because all point forecasts revert fairly quickly to the climatological forecast, and hence all long-horizon forecasts are "equally poor."

On reflection, however, our model's point forecasting performance is also not particularly discouraging, insofar as the crucial forecasts for weather derivatives are not point forecasts, but rather density forecasts. That is, a key object in any statistical analysis involving weather derivatives—indeed, the key object for the central issue weather derivative pricing—is the entire conditional density of the future weather outcome. The point forecast is the conditional mean, which describes just one feature of that conditional density, namely its location. Hence the fact that the long-horizon conditional mean estimate produced by our model is no better than that produced by the climatological or EarthSat models does not imply that our model or framework fails to add value. On the contrary, a great virtue of our approach is its immediate and simple generalization to provide entire density forecasts via stochastic simulation. In particular, the main feature of average temperature conditional density dynamics, apart from the seasonal conditional mean dynamics, is the highly seasonal conditional variance dynamics, which we have modeled parsimoniously and successfully. This facilitates simple modeling of time-varying scale of the conditional density, and *it is as relevant for very long horizons as for very short horizons.*

All of this adds up to a simple, yet potentially powerful framework for producing density forecasts of weather variables, to which we now turn. It is telling to observe that in what follows we must evaluate the performance of our density forecasts in absolute terms, rather than relative to EarthSat density forecasts, because EarthSat, like most weather forecasters, does not produce density forecasts.

3.2 Density Forecasting

In this section we shift our focus to long-horizon density forecasting and to cumulative heating degree days, all of which is of crucial relevance for weather derivatives. Heating degree days for day t is simply $HDD_t = \max(0, 65 - T_t)$. We use our model of daily average temperature to produce density forecasts of cumulative HDDs (*Cum HDDs*) from November 1

through March 31, for each city and for each year between 1960 and 2000, defined as $CumHDD_{y,i} = \sum_{t=1}^{151} HDD_{t,y,i}$, for $y = 1960, \dots, 2000$, $i = 1, \dots, 4$. Because we remove February 29 from each leap year, each sum contains exactly 151 days. We use full-sample as opposed to recursive parameter estimates, as required by the very small number of CumHDD observations. To avoid unnecessarily burdensome notation, we often drop the y and i subscripts when the meaning is clear from context.

We focus on CumHDD for two important reasons. First, weather derivative contracts are often written on the cumulative sum of a weather-related outcome over a fixed horizon, as with the cumulative HDD and CDD contracts traded on the CME. Second, and related, the November–March HDD contract is one of the most actively traded weather-related contracts and hence is of substantial direct interest.

On October 31 of each year, and for each city, we use the estimated daily model to produce a density forecast of CumHDD for the following winter’s heating season. We simulate 250 realizations of CumHDD, which we then use to estimate the density, as follows. First, we simulate 250 151-day realizations of the temperature shock, ε_t , by drawing with replacement from the empirical distribution of estimated temperature shocks, $\hat{\varepsilon}_t$. Second, we run the 250 151-day realizations of temperature shocks through the estimated model (1) to obtain 250 simulated 151-day realizations of daily average temperature. Third, we convert the 250 simulated 151-day realizations of daily average temperature into 250 simulated 151-day realizations of HDD, which we cumulate over the November–March heating season, $CumHDD_s = \sum_{t=1}^{151} HDD_{t,s}$, $s = 1, 2, \dots, 250$. Finally, we form the empirical distribution function of CumHDD based on $CumHDD_s$, $s = 1, \dots, 250$.

After passing through the entire sample, we have 41 assessed distribution functions, $\hat{F}_y(\cdot)$, $y = 1960, \dots, 2000$, with 1 function governing each of $CumHDD_y$, $y = 1960, \dots, 2000$. We assess the conditional calibration of those distributional forecasts via the probability integral transform, as suggested by Rosenblatt (1952) and extended by Dawid (1984), Diebold, Gunther, and Tay (1998), and Diebold, Hahn, and Tay (1999). In particular, if the estimated distribution and true distribution coincide year by year, then the series of assessed distribution functions, $\hat{F}_y(\cdot)$, evaluated at the corresponding series of realized values of $CumHDD_y$ should be approximately iid and uniformly distributed on the unit interval. Formally, $z_y \equiv \hat{F}_y(CumHDD_y) \stackrel{iid}{\sim} U(0, 1)$. For each city, we check uniformity by examining histograms of z and check independence by examining correlograms of the first four powers of z . The sample of size 41 is of course small, but the framework has been previously applied successfully in small samples, as by, for example, Diebold, Tay, and Wallis (1999).

First, consider assessing uniformity. We estimate the density of z using simple four-bin histograms, presented in the left-most column of Figure 9, accompanied by 95% pointwise error bands under the iid $U(0, 1)$ null hypothesis. Interestingly, the z series differ rather noticeably from uniformity, and moreover they display a common pattern; too many large CumHDD realizations occur relative to the assessed distributions, as evidenced by the increase in the histograms when moving from left to right. The common nature of uniformity violations may

indicate a neglected common temperature component, due to, for example, El Niño, La Niña, changes in the jet stream, or various other global factors.

Now consider assessing independence. The last four columns of Figure 9 show the correlograms of the first four powers of z , taken to a maximum displacement of 10 years, together with asymptotic 95% Bartlett bands under the iid null hypothesis. The results are mixed, but a common pattern of some positive serial correlation is often apparent.

We view our CumHDD distributional forecasting performance as encouraging, although there is clear room for improvement. Evidently the effects of small specification errors in the daily model, which have negligible consequences for near-term forecasting, cumulate as the horizon lengthens, producing large consequences for longer-term forecasting. The error in forecasting CumHDD is of course the sum of the many component daily errors, and the variance of that sum is the sum of the variances *plus* the sum of all possible pairwise covariances. Hence tiny and hard-to-detect but slowly-decaying serial correlation in 1-day-ahead daily average temperature forecasting errors may cumulate over long horizons. In future work beyond the scope of this article, it will be of interest to attempt to address the specification error issue by modeling and forecasting CumHDD directly. In contrast, currently we fit only a single (daily) average temperature model, which we estimate by minimizing a loss function corresponding to 1-day-ahead mean squared prediction error, then use the model to produce forecasts at many different horizons, all of which feed into our CumHDD forecasts.

4. CONCLUDING REMARKS AND DIRECTIONS FOR FUTURE RESEARCH

Weather modeling and forecasting are crucial to both the demand side and the supply side of the weather derivatives market. On the demand side, to assess the potential for hedging against weather surprises and formulate the appropriate hedging strategies, one needs to determine how much “weather noise” exists for weather derivatives to eliminate. This requires weather modeling and forecasting. On the supply side, standard approaches to arbitrage-free pricing are irrelevant in weather derivative contexts, and so the only way to price options reliably is again by modeling and forecasting the underlying weather variable. Rather curiously, it seems that little thought has been given to the crucial question of how best to approach weather modeling and forecasting in the context of weather derivative demand and supply. The vast majority of extant weather forecasting literature has a structural “atmospheric science” feel, and although such an approach is surely best for forecasting at very short horizons, as verified both by our own results and those of many others, it is not obvious that it is best for the longer horizons relevant for weather derivatives, such as 12 weeks or 6 months. Moreover, density forecasts, but not point forecasts, are of maximal relevance in the derivatives context. Good distributional forecasting does not necessarily require a structural model, but it does require accurate approximations to stochastic dynamics.

In this article we took an arguably naive nonstructural time series approach to modeling and forecasting daily average temperature in four U.S. cities, and we inquired systematically as

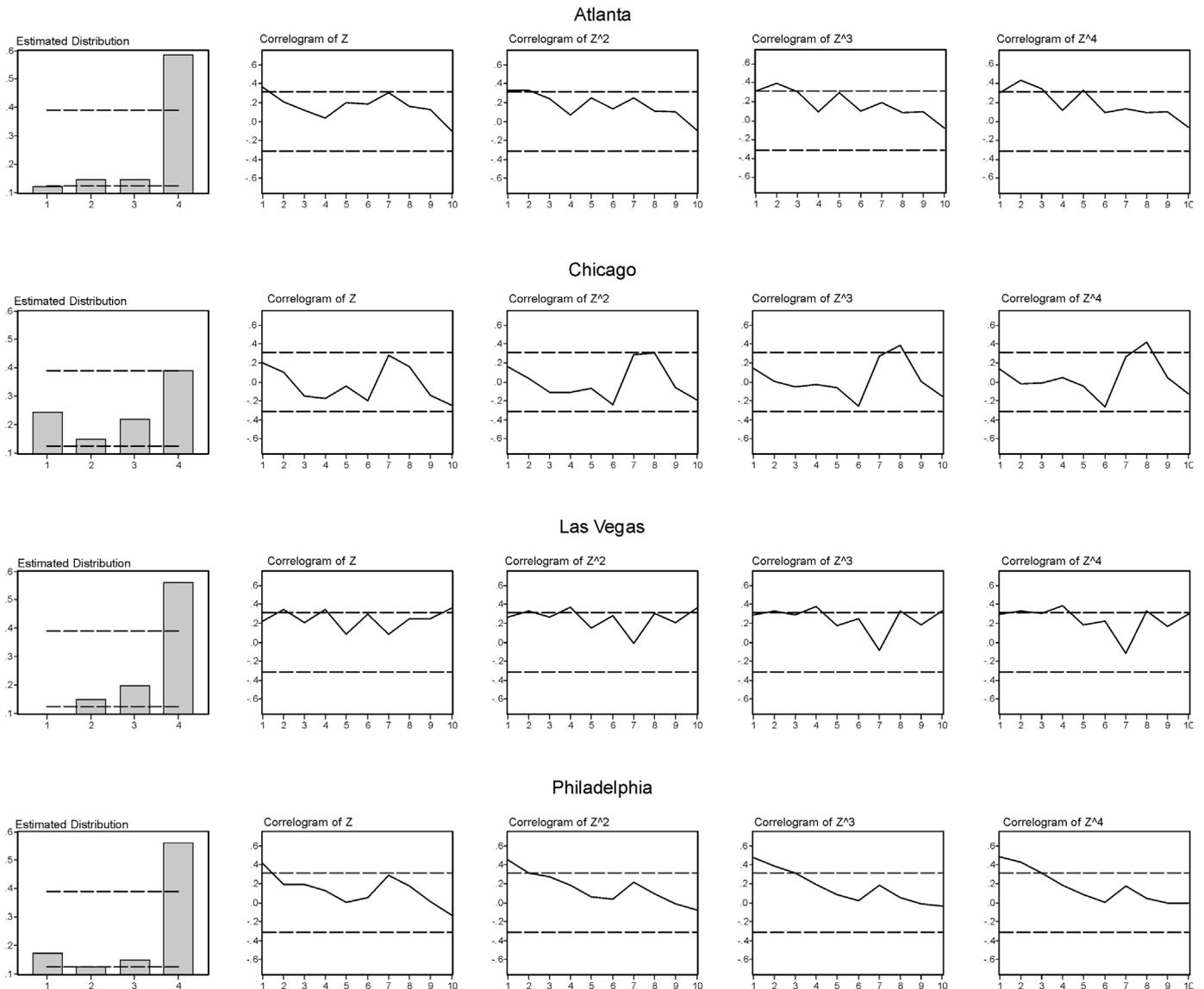


Figure 9. *z*-Statistics Distributions and Dynamics, Daily Average Temperature Distributional Forecasts. Each row displays a histogram for *z* and correlograms for four powers of *z*, the probability integral transform of cumulative November–March HDDs, 1960–2000. Dashed lines indicate approximate 95% confidence intervals in the iid $U(0, 1)$ case of correct conditional calibration.

to whether it proves useful. The answer, perhaps surprisingly, was a qualified yes. Our point forecasts were always at least as good as the persistence and climatological forecasts, but were still not as good as the judgementally adjusted NWP forecast produced by EarthSat until a horizon of 8 days, after which all point forecasts performed equally well. Crucially, we also documented and modeled strong seasonality in weather surprise volatility, and we assessed the adequacy of long-horizon distributional forecasts that accounted for it, with mixed but encouraging results. Moreover, we found an interesting commonality in the patterns of cross-city deviations from perfect conditional calibration, indicating possible dependence on common latent components, due perhaps to El Niño or La Niña.

The key insight is that the losses associated with the non-structural approach, which bypasses atmospheric data and science in favor of statistical extrapolation, although surely important for very short-term forecasting, may be largely irrelevant when forecasting several months ahead, as is typically

required for weather derivatives. Instead, time and money may be better spent developing simple statistical models useful for density forecasting, because it appears that simple, yet sophisticated time series models and forecasts perform at least well enough to suggest the desirability of additional exploration. The time series models are:

1. Parsimonious and simple. Only a few parameters need to be estimated, and only standard statistical methods are used.
2. Flexible. The model can capture dynamics that may or may not contain trend, seasonality, and cycles.
3. Extensible. The model may readily be modified to accommodate additional features if desired—even “structural” features related for example to occurrence of El Niño or La Niña.
4. Inexpensive. Analysis and forecasting with the model requires only standard, widely available, and inexpensive software and data and minimal human maintenance and

oversight, facilitating not only model and forecast construction, but also replication of results.

5. Intrinsically stochastic and focused on the entire conditional distribution, not just on the conditional mean. Hence the approach is naturally suited to the construction and interpretation of the long-horizon climatological forecasts, *stated as complete densities*, as needed in weather derivatives contexts.

Hence we believe that a strong case exists for their use in the context of modeling and forecasting as relevant for weather derivatives.

We would also assert that our views are consistent with the mainstream consensus in atmospheric science. For example, in his well-known text, Wilks (1995, p. 159) noted that “Statistical weather forecasting methods” are still viable and useful at very short lead times (hours in advance) or very long lead times (weeks or more in advance) for which NWP information is either not available with sufficient promptness or accuracy, respectively.” Indeed, in many respects our results are simply an extensive confirmation of Wilks’ assertion in the context of weather derivatives, which are of great current interest.

Ultimately, our present view on weather forecasting for weather derivatives is that climatological forecasts are what we need, but that traditional point climatological forecasts—effectively little more than daily averages—are much too restrictive. Instead, we seek “generalized climatological forecasts” from richer models tracking entire conditional distributions, and modern time-series statistical methods may have much to contribute. We view this article as a “call to action,” with our simple model representing a step toward a fully generalized climatological forecast, but with many important issues remaining unexplored. Here we briefly discuss a few that we find particularly intriguing.

One of the contributions of this article is our precise quantification of daily average temperature-conditional variance dynamics. But richer dynamics might be beneficially permitted in both lower-ordered conditional moments (i.e., the conditional mean) and higher-ordered conditional moments (e.g., the conditional skewness and kurtosis). In regard to the conditional mean, we could introduce explanatory variables, as was done by Visser and Molenaar (1995), who condition on a volcanic activity index, sunspot numbers, and a southern oscillation index. Relevant work also includes that of Jones (1996) and Pozo et al. (1998), but those authors used annual data and thus missed the seasonal patterns in both conditional mean and conditional variance dynamics so crucial for weather derivatives demand and supply. We could also allow for nonlinear effects, most notably stochastic regime switching in the tradition of Hamilton (1989), which might aid in, for example, the detection of El Niño and La Niña events (see Richman and Montroy 1996; Zwiers and von Storch 1990). In terms of the conditional skewness and kurtosis, we could model them directly, as for example, with the autoregressive conditional skewness model of Harvey and Siddique (1999). Alternatively, we could directly model the evolution of the entire conditional density, as was done by Hansen (1994).

Aspects of multivariate analysis and cross-hedging also hold promise for future work. Cross-city correlations may be crucially important, because they govern the potential for cross-hedging. Hedging weather risk in a remote Midwestern location

might, for example, be prohibitively expensive or even impossible due to illiquid or nonexistent markets, but if that risk is highly correlated with Chicago’s weather risk, for which a liquid market exists, then effective hedging may still be possible. Hence an obvious and important extension of the univariate temperature analysis reported in this article is multivariate modeling of daily average temperature in a set of cities, allowing for a time-varying innovation variance–covariance matrix. Of particular interest would be the fitted and forecasted conditional mean, conditional variance, and conditional covariance dynamics; the covariance matrices of standardized innovations; and the impulse response functions (which chart the speed and pattern with which weather surprises in one location are transmitted to other locations).

Another interesting multivariate issue involves weather-related swings in earnings and share prices. It will be of interest to use the size of weather-related swings in earnings as a way to assess the potential for weather derivatives use. In particular, we need to understand how weather surprises translate into earnings surprises, which then translate into stock price movements. Some interesting subtleties may arise. As one example, note that only systematic weather risk should be priced, which raises the issue of how to disentangle systematic and nonsystematic weather risks. As a second example, note that there may be nonlinearities in the relationship between prices and the weather induced via path dependence; for example, if there is an early freeze, then it does not matter how good the weather is subsequently; the crop will be ruined, and prices will be high (see Richardson, Bodoukh, Sjen, and Whitelaw 2001).

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